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THESIS ABSTRACT

Name: Richard S. Cheney

Title: Geophysical Investigation of the Raton Basin

Rank: Captain, USAF, 1982

Degree: Master of Science in Geosciences, Texas Tech University

This thesis correlates gravity, magnetic, and seismic data for the Raton Basin of Colorado and New Mexico. The gravity data suggest that the study area, and the region around it, is in isostatic equilibrium. The free air anomaly in the southern portion of the study area suggests lack of local compensation due to Quaternary volcanic rock. The volcanic rock thickness, calculated from the free air gravity data, is 180 m. The gravity data indicated a crustal thickness of about 45 km, and the crust thinned from west to east.

A basement relief map was constructed from the Bouguer gravity data. Computer techniques were developed to calculate the depth to the basement surface and to plot a contour map of that surface. The Raton Basin magnetic map defined the same surface found on the basement relief map since the overlying sedimentary rocks have no magnetism; therefore, any magnetism present is caused by the basement rock.

A seismic survey near Capulin Mountain detected a high level of microseismicity that may be caused by adjustment along faults or dormant volcanic activity.



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GEOPHYSICAL INVESTIGATION OF THE RATON BASIN

bу

RICHARD STEPHEN CHENEY, A.B., M.A.

A THESIS

IN

GEOSCIENCE

Submitted to the Graduate Faculty of Texas Tech University in Partial Fulfillment of the Requirements for the Degree of

MASTER OF SCIENCE

Approved

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CHAPTER I

INTRODUCTION

General Statement

The Raton Basin of New Mexico and Colorado (35°-38°N; 104°-105.5°W) is a structural depression which is defined by the Sangre de Cristo Up-lift on the West, the Wet Mountains Uplift to the north, the Apishipa Arch on the northeast, and the Sierra Grande Arch on the southeast (figure 1). Today the Raton Basin is outlined by topographic relief caused by the caprock "effect" of the Trinidad Sandstone. The importance of the outlining "effect" of the Trinidad Sandstone is illustrated by both Baltz (1965) and Brill (1952) who draw the eastern boundary of the Raton Basin at the outcrop of Trinidad Sandstone, rather than along the axes of the Apishipa and Sierra Grande Arches.

Extending 175 miles long and a maximum of 65 miles wide, the Raton Basin is divided by the Cimarron Arch into the northern Raton Basin and the Las Vegas Basin on the south. The deepest depression within the Basin is located near the Sangre de Cristo Uplift. Rocks on the west limb of the Basin dip vertically, whereas rocks on the east limb dip one to five degrees (Wanek, 1963). In this thesis, the entire Raton Basin will be referred to as the Basin, while the northern portion will be called the Raton Basin, and the southern portion referred to as the Las Vegas Basin.

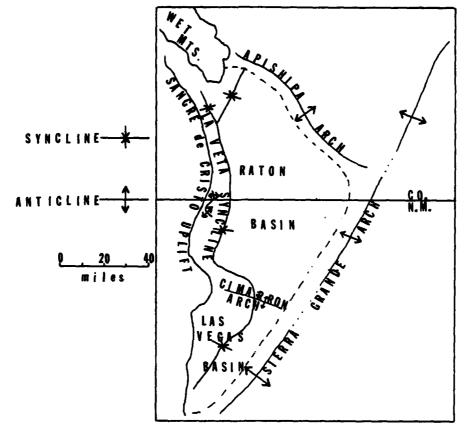


Figure 1. Geologic structures defining the Raton Basin of New Mexico and Colorado.

Purpose of the Investigation

Knowledge of the Basin is incomplete, and correlation of geophysical and geologic data has not been previously published. Therefore, in order to fill this void in the literature this thesis will undertake the following objectives: 1) Obtain and interpret both Free Air and Bouguer Gravity maps; 2) Obtain and interpret a vertical component magnetic map; 3) Obtain vertical, short-period seismic records for the Capulin area in an effort to determine the natural seismic activity; 4) Compare

geophysical data with the geologic structure, and 5) Determine whether economic potential in the Basin can be inferred from geophysical data. The study being reported here includes gravity, magnetic, and seismic surveys of portions of the Basin, including the Sierra Grande Arch and adjacent High Plains (figure 2).

General Geology of the Basin

The first of the two orogenies that have affected the Basin occurred during Pennsylvanian and Permian time (Baltz, 1965), and was responsible for detritus filling the Central Colorado Basin (located in the northern portion of the Raton Basin). The detrital material was derived from both west (San Luis Uplift) and east (the ancestral Wet Mountains Uplift) of the Basin. Contemporaneously, the Rowe-Mora Basin, which was in the general vicinity of the present Las Vegas Basin, also formed. This Basin was defined by the San Luis Uplift to the west, the Sierra Grande Uplift to the east, the Cimarron Arch on the north, and was linked by a low saddle to the Tucumcari Basin to the south. The Colorado and Rowe-Mora Basins filled with sediments which were derived from the ancestral Rockies and deposited in a marine transgressive environment of Pennsylvanian age. The resulting strata show a laterally complex composition reflecting the mixed terrestial-marine origin. These strata are thickest in the regime of the Sangre de Cristo Mountains. Large accumulations of sandstone and limestone occur in the southern portion of the Rowe-Mora Basin, but shale deposits, which are dominated by thick grey to black strata, predominate to the north. In fact, this shale which

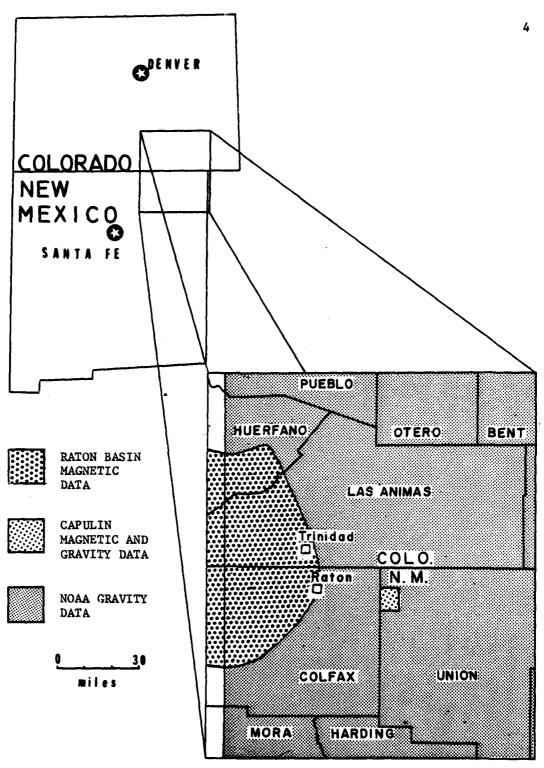


Figure 2. Location of the study area and limits to each survey in this report.

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characterizes the northern part of the ancient Central Colorado Basin is a minimum of 9,500 feet thick (Brill, 1952). On the margins of the ancient basins, shale deposits thin towards the San Luis, Sierra Grande, and Wet Mountains Uplifts.

The depositional history of the strata in the Basin is complex as is the correlation between contemporaneous formations in the Las Vegas and Raton Basins (figure 3). Strata which were deposited in the Permian era are dominated by marine shales and clastic rocks. Triassic strata are composed primarily of terrestial clastic rocks but are, however, sparsely represented in the area. Today the Las Vegas Basin area is covered partially by Quaternary volcanic rocks. Discussion of the depositional and orogenic history is limited to the area of the present Basin and is based on rock outcrops and the known geologic structure (plate 1).

The oldest rocks identified in the Basin are marine sandstones and limestones of the Devonian Espiritu Santo Formation (Baltz, 1965) in the Las Vegas Basin and the Devonian Chaffee Formation (Johnson, 1969) in the Raton Basin. Limestone breccia of the Mississippian Tererro Formation rests unconformably on the Espiritu Santo Formation in the Las Vegas Basin. The Tererro Formation is missing in the Raton Basin.

The Pennsylvanian Magdalena Group is composed of sandstone and grey shale of the Sandia Formation and limestone of the overlying Madera Formation. The Sandia Formation in the Las Vegas Basin, which consists of over 1,000 feet of sedimentary rock, is correlative with the Deer Creek Formation and the lower part of the Minturn Formation of the Raton Basin.

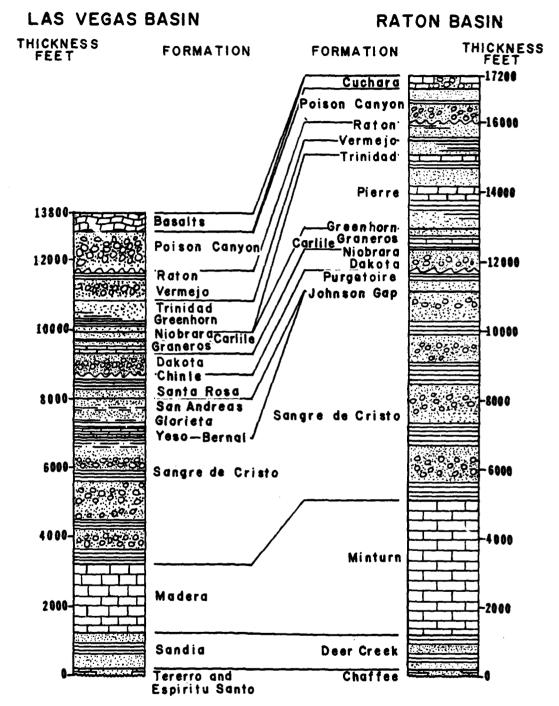


Figure 3. Columnar section of the Las Vegas and Raton Basins. Formation thicknesses are from Baltz (1962); symbols used in the columns are from Compton (1962).

The Madera limestone ranges from 1,000 to 3,000 feet in thickness in the Las Vegas Basin; it grades laterally northward into the Minturn Formation, which has a thickness of 5,500 feet. These units thin to the west and east of the Basin in the general direction of the San Luis and Sierra Grande Uplifts.

The Upper Pennsylvanian/Permian Sangre de Cristo Formation is approximately 3,500 feet thick in the Las Vegas Basin and greater than 9,500 feet thick in the Raton Basin (figure 3).

At the end of the Permian time a sea filled the Rowe-Mora Basin, and deposition of siltstone and shale of the Yeso Formation began.

Located above the Yeso Formation is the Glorieta Sandstone and San Andreas Limestone, respectively. Overlying the San Andreas Limestone is the Bernal Formation which is composed of limestone, siltstone, shale, and sandstone. These three formations have a combined, maximum thickness of 1,075 feet (figure 3).

The Late Triassic Dockum Group lies conformably on Permian strata. In the Las Vegas Basin the Group is composed of both the Santa Rosa Sandstone and the overlying Chinle Formation shales. North of the Cimarron Arch the Dockum Group is correlative to the Johnson Gap Formation, which is composed of limestone conglomerate, limestone, siltstone, shale, and sandstone (Baltz, 1965). The Dockum Group is approximately 1,100 feet thick in the Las Vegas Basin and less than 700 feet thick in the Raton Basin.

By Early Cretaceous time the region was part of the Rocky Mountain geosyncline. The Laramide Orogeny, the second orogeny to affect the

area, continued from Cretaceous time through Oligocene time and produced the present Las Vegas and Raton Basins and other associated structures. At the beginning of the orogeny, the San Luis Uplift was still a positive area. By Eocene time the present Raton Basin was defined by the marked uplift of the Sangre de Cristo region and the gentle uplift of the Apishipa-Sierra Grande Arches. The Sangre de Cristo Uplift merged with the San Luis Uplift prior to the formation of the Rio Grande rift. When the Rio Grande rift formed in Miocene/Pliocene time, the San Luis positive area was downfaulted into the rift graben.

The Early Cretaceous Purgatoire Formation is composed of both conglomeratic sandstone and sandstone. The Dakota Sandstone overlies the Purgatoire Formation and is composed of interbedded, buff sandstone and shale. Together, these two formations are up to 650 feet thick, and today they form the primary groundwater aquifer in the Basin. Estimates place the volume of water stored in the Dakota Sandstone at 55 million acre-feet in New Mexico (Griggs, 1948); twice the above volume is estimated for the Basin.

The Graneros Shale, Greenhorn Limestone, and Carlile Shale are Late Cretaceous in age and vary in thickness from 385 feet to 700 feet (Baltz, 1965). The Upper Cretaceous Niobrara Formation is composed of less than 100 feet of shale and sandstone and rests conformably on the Carlile Shale.

Located on top of the Niobrara Formation is the Upper Cretaceous Pierre Shale, which is up to 2,300 feet thick in the Raton Basin but is absent in the Las Vegas Basin. The formation is composed of shale,

thin beds of limestone, and sandstone. Above the Pierre Shale are the Upper Cretaceous Trinidad Sandstone and Vermejo Formation, which are undifferentiated on some geological maps (Bachman and Dane, 1962). The Trinidad Sandstone is composed of arkosic sandstone with thin interbedded shale; it intertongues with the shale, coal, and arkosic sandstone of the Vermejo Formation. The thickness of the two formations varies from 250 feet to 850 feet. The Vermejo Formation is the main source of coal in the Basin, producing high volatile C bituminous coal of coking quality (Wanek, 1963). In the areas of extensive intrusive rocks, much of the coal was destroyed by contact metamorphism (Jurie and Gerhard, 1969).

The Raton Formation is Late Cretaceous to Early Tertiary in age and overlies the Vermejo Formation. It has a maximum thickness of 1,700 feet and is composed of arkosic sandstone, shale, and coal. The Poison Canyon Formation of Paleocene age lies unconformably on the Raton Formation. It consists of as much as 2,500 feet of arkosic sandstone, conglomerate, and thin shale. The youngest significant formation in the Basin is the Cuchara Formation, which is exposed around Spanish Peaks. It is Eocene in age and is composed of conglomeratic sandstone and interbedded shale.

During the latter stages of the Laramide Orogeny numerous intrusive rocks penetrated the sedimentary rocks. Dike swarms, sills, stocks and laccoliths dating from Eocene/Oligocene time are found throughout the Basin. However, the greatest concentration is found in the Spanish Peaks intrusive area. Many small anticlines are attributed to the intrusive

upwelling. Emplacement of many of the dikes, which are silicic to ultra basic in composition, appears to have been fault controlled (Wanek, 1963).

Volcanic activity was concentrated near the Sierra Grande Arch. Baldwin and Muehlberger (1959) identified three sequences of volcanic flows, dating from Miocene to Recent, that they called Raton, Clayton, and Capulin, respectively. Both the Raton and Capulin sequences are composed of basalts, and the Clayton sequence has a varied composition which ranges from extremely alkalic to subsilicic. Geochemical analysis of the volcanic rock suggests the magma originated in the upper mantle (Jones et al., 1976). Stormer (1972) suggests eastward migration of the vulcanism across the Rio Grande rift, from the San Juan volcanic field to the Sierra Grande volcanic field, occurred. Sanford et al., (1981) also suggest that the Sierra Grande vulcanism might be an eastern extension of the Jemez Lineation of western New Mexico. Further, Hamilton and Pakiser (1965) indicate that the Rio Grande rift is limited to the upper crust. Thus, the mantle may be the magma source in vulcanism on both sides of the rift.

It has been suggested (Edwards et al., 1978) that anomalously high heat flows may be caused by igneous activity within the Basin. Typical heat flow values for the Great Plains average 1.5 HFU (1 HFU = lµcal/cm² sec), but values as high as 4.7 HFU are found in the Basin. Suppe et al., (1975) proposed a hot spot trace from eastern Arizona to a location near Raton, New Mexico. Reiter et al., (1979) theorized that a thermal anomaly formerly existed east of the Southern Rocky Mountains and is now

centered near the Sierra Grande Arch. They noted that radioactivity of the sedimentary rocks within the Basin can account for only ten percent of the observed increased heat flow; the remainder is unexplained.

In summary, the total thickness of sedimentary fill in the Basin is about 17,000 feet. The Basin sediments date from Middle Paleozoic time and have been affected by two orogenies. These sedimentary rocks have economic importance, and mining activities in the Raton Basin have produced coal and graphite in commercial quantities. Cretaceous rocks are also known to have both oil and gas in small quantities. Currently, there is one commercial oil well in Huerfano County, Colorado, at the northern end of the Basin, and a gas field near the Basin in eastern Las Animas County, Colorado. A basement structure map will be developed in Chapter III using the densities of the sediments described above, and the gravity data that will be introduced in Chapter II.

CHAPTER II

DATA COLLECTION AND REDUCTION

General Statement

A portion of the gravity data used in this study were obtained from the National Oceanic and Atmospheric Administration (NOAA), Washington, D.C. The gravity and magnetic measurements in the vicinity of Capulin National Monument, New Mexico and the magnetic measurements in the Raton Basin were made by the author during 1981. Each gravity station in the vicinity of Capulin National Monument was also a magnetic station. In order to provide a detailed analysis of the Basin, additional magnetic stations were spaced midway between the gravity stations. NOAA station density is 0.06341 station per square mile; Capulin gravity station density is 1.333 stations per square mile; and Raton Basin magnetic station density is 0.0650 station per square mile.

Elevation Control

U.S. Geological Survey fifteen minute quadrangle maps were used to determine elevations for stations near Capulin. Such elevation data were supplemented with a preliminary Bureau of Land Management map of Capulin National Monument. Spot elevations on the maps at road intersections and benchmarks were used wherever possible. Elevation accuracy is within five feet, corresponding to ± 0.35 mgal accuracy in gravity measurement. Elevation effects were insignificant in the reduction of

the magnetic data.

Magnetic Data Measurement and Reduction

The Capulin magnetic data were collected using an E.J. Sharpe PMF-3 flux gate magnetometer (50 and 100 gammas per division). Station number one (figure 14) was arbitrarily chosen by the author as the datum for the survey. Station reoccupation resulted in a minimum of three datum measurements each day and a minimum of two cross correlations on stations occupied on previous days. Diurnal correction graphs were computed daily, and were used to adjust all measurements.

The Raton Basin magnetic data were collected using an Askania vertical Torsion magnetometer (2.433 gammas/degree-scale division). Field procedures for the Raton magnetic survey were the same as those employed in the Capulin magnetic survey. Normal field corrections were computed from 36°30'N, 105°15'W using the Regional Vertical Intensity Map 1965. Latitude correction was 909.09 gammas per degree latitude (12.98 gammas per mile); longitude correction was 434.80 gammas per degree longitude (7.9 gammas per mile).

Gravity Data Measurement and Reduction

Datum for gravity data acquired from NOAA is sea level. The density used in Bouguer anomaly calculations was 2.67 g/cc. Gravity data acquired near Capulin National Monument were collected using a North American gravimeter (0.107 mgal per division). Station number one was

selected to coincide with a NOAA station in the survey area and was used as the datum for the Capulin survey. This procedure allowed comparison of the Capulin data to the NOAA data. Latitude corrections were made from 36°48'18"N at a rate of 1.25 mgal per mile.

Seismic Data

A short period vertical Geotech 18300 seismometer with a free period of one second was buried approximately two feet in depth at the Visitor's Center, Capulin National Monument, from 16 July 1981, to 23 September 1981. During this period, 51 daily seismograms containing 1,122 hours of data were recorded by a Sprengnether VR-50 recorder-amplifier. Damping was increased after 14 August 1981 from normal internal resistance to 2,200 ohms resistance across the input terminals of the amplifier. Filters were not used.

Each seismogram contains twenty-four hours of data with the minute marks spaced sixty millimeters apart. The ink trace was approximately 1.0 mm thick and makes it impossible to resolve frequencies higher than one Hertz. Estimated amplification from artificial inputs is between 25,000 to 50,000.

Computer Processing Techniques

Modified computer programs from Davis (1973) and new programs written by the author for this study were employed to process both magnetic and gravity data on an IBM 360 computer housed at Texas Tech University. The machine processing helped to verify manual calculations and also to derive mathematical models that would have been impossible to obtain by

other means.

All data were systematically processed in the same manner. First, the corrected data were filtered by a uniform matrix of either 1,600 or 2,400 elements. The matrix value of each element was obtained by calculating the algebraic average of the six nearest data values. Development of the matrix by interpolation over six data values for each element results in the minimum and the maximum values in the completed matrix to be less than the minimum and the maximum values in the input data. Less than six data values in the interpolation increased the minimum-maximum range in the matrix; however, the areal extent of the maximum and the minimum features would have appeared exaggerated in the final map. This causes mountain peaks to appear as elevated buttes or mesas. The use of more than s < data values in interpolation resulted in excessively reducing the minimum-maximum range in the matrix.

Statistical Analysis System's SAS/GRAPH (Reinhardt, 1981) was used to produce the computer maps from the interpolated matrix. These maps were contoured on a strict mathematical ratio by the computer; therefore, the technique can be used advantageously to compare different data sets (gravity and magnetic) produced at the same scale and in the same geographic area. In order to compare the different mapping techniques, both computer plotted and hand drawn maps using the same data sets have been included in this report.

The computer also was used to compare the computer contoured map of the data contained in matrix form to a surface defined by a polynomial of order n. A polynomial of order O represents a dipping plane, and polynomials of very high order very accurately represent the data in matrix form. A second order polynomial was found to accurately represent the steep gravity dip under the Sangre de Cristo Mountains and the shallow gravity dip of the Great Plains. The residual gravity (Smith, 1970; Dobrin, 1976), or difference, of the second order polynomial from the matrix gravity data was used to construct the basement structure map described in Chapter III.

CHAPTER III

PRESENTATION AND INTERPRETATION OF THE GRAVITY, MAGNETIC, AND SEISMIC DATA

Free Air Gravity Data

The NOAA free air gravity data have been tabulated and included in Appendix A. The gravity data acquired in the Capulin area as part of this study, which will be discussed as an addition to the NOAA data, have been tabulated and included in Appendix B.

The NOAA free air gravity anomaly data are shown, for comparison, in hand drawn (figure 4) and computer plotted (figure 5) formats. The algebraic averages for the station elevation, free air anomaly, and Bouguer anomaly in each one degree block in the study area are shown in figure 6. The free air anomalies for the northern two blocks of the study area are nearly zero and suggest isostatic equilibrium for this area. However, the free air anomalies are significantly positive in the southern two blocks of the study area and could indicate lack of local isostatic adjustment in these areas. The question of regional isostatic equilibrium is, therefore, raised.

Perfect isostatic adjustment assumes that any near surface positive mass has an equal amount of negative mass below the compensation level within the Earth, and in the simple case, such a mass distribution would cause near zero free air gravity anomalies. However, in elevated terrain, as in the study area, the gravitational effect upon a surface

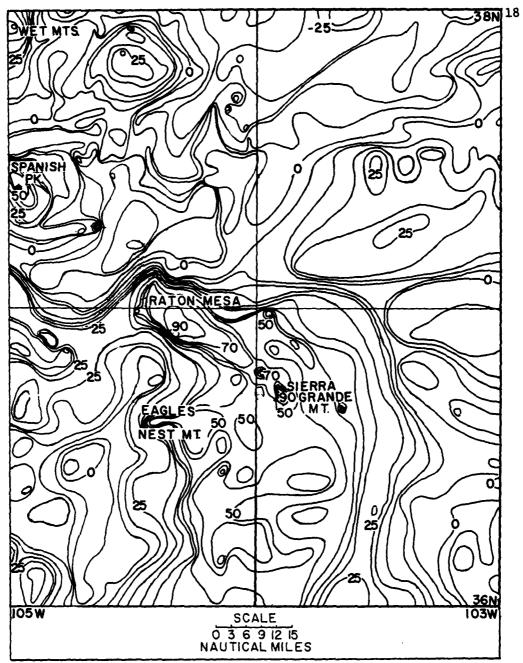


Figure 4. Regional free air gravity map using NOAA gravity data.

Datum is sea level and gravity stations are shown in

Figure 10. Contoured by hand.

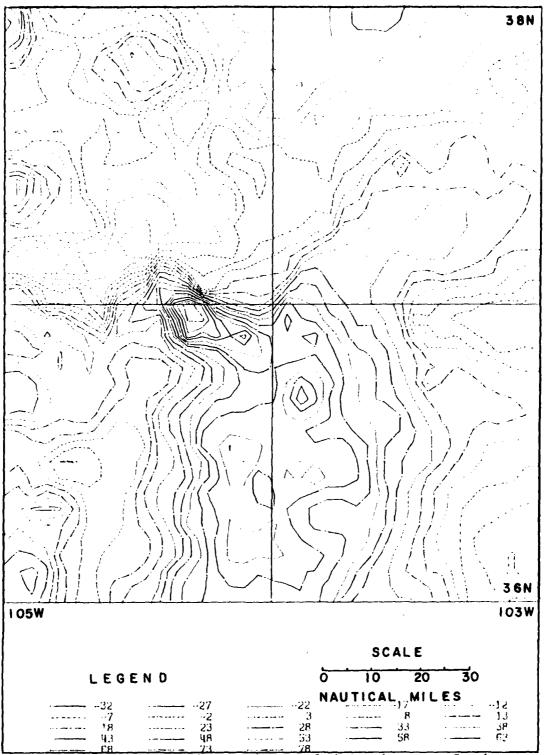


Figure 5. Regional Free Air anomaly map of the study area. Gravity stations are the same as in figure 10. Contour interval is 5 mgal, datum is sea level.

	38N
E = 1818 m FA = -1.28 MGAL BA = -204.7 MGAL EE = 1832 m FC = +1.6 MGAL CFA = +0.4 MGAL	E = 1412 m FA = -0.90 MGAL BA = -158.8 MGAL EE = 1421 m FC = +1.0 MGAL CFA = +0.1 MGAL
E = 2010 m FA = +26.2 MGAL BA = -198.5 MGAL EE = 1776 m FC = -26.1 MGAL CFA = +0.1 MGAL	E = 1636 m FA = +19.9 MGAL BA = -163.4 MGAL EE = 1462 m FC = -19.4 MGAL CFA = +0.5 MGAL
105 W	1 0 3 W

Figure 6. Graphic representation of the algebraic averages of the NOAA gravity and elevation data divided into one degree latitude and longitude blocks. E = elevation; FA = free air anomaly; BA = Bouguer anomaly; FC = Faye correction; EE = equivalent elevation; CFA = corrected free air anomaly (FA + FC).

measurement, of the near surface rock (the positive mass) is larger than the gravitational effect of the negative mass at the compensation level simply because the near surface rock is nearer the measured station. This means that a positive free air anomaly is in part caused by depth to compensation and does not necessarily indicate crustal loading. It is possible that the total observed free air anomaly can be explained in terms of the depth of isostatic compensation. Therefore, if the free air anomaly is to be used in an estimation of isostatic adjustment, a correction must be applied.

The correction applied to the observed free air anomaly for the effect of depth to compensation is generally called the Faye correction, after the man who first discussed the effect in the Nineteenth Century. One method of calculating the Faye correction uses the Bouguer anomaly to determine the elevation of the surface of a slab which is in isostatic balance (Woollard, 1962). This is done by assuming if isostatic adjustment is complete, the Bouguer anomaly is related to the surface elevation of the crustal load causing the adjustment. The elevation of this surface slab is thought of as the equivalent elevation (h') derived from the Bouguer anomaly (h' = $BA/2\pi\gamma\rho$; where γ is the gravitational constant and p is the assumed density). The equivalent elevation is subtracted from the station elevation (h) and the difference (Δh) defines the slab thickness used in the Faye correction (F.C.): F.C. = $2\pi\gamma\rho\Delta h$. The Faye correction is subtracted from the free air anomaly to produce a corrected free air anomaly value, which is actually the isostatic anomaly since this computation has removed the effect of depth of compensation upon

the free air anomaly. As a result of applying the Faye correction to the free air anomalies in the study area (figure 6) the corrected free air anomalies clearly suggest regional isostatic equilibrium.

Although regional isostatic equilibrium is established, the locally positive average free air anomalies of the southern two blocks (figure 6) are indications of a lack of local compensation. The Quaternary volcanic rock covering most of the southern two blocks of the study area is missing in the northern portion of the study area and is assumed to be the cause of the large positive free air anomaly in the southern two blocks. If this is so, the free air anomaly can be used to estimate the rock thickness. If a density of 2.67 g/cc is assumed for the volcanic rock, and the formula for the gravitational effect of a slab is solved, a thickness of 180 m is indicated for the Quaternary volcanic rocks. As a confirmation of the rock thickness calculated from gravity data, a water well at Capulin National Monument penetrated 196 m of volcanic rock. The real thickness is probably somewhat irregular, and the value computed from the gravity data appears to be a good estimate.

Comparison of the free air map and topographic map (figure 7) shows that positive topographic features are also areas of positive free air gravity anomalies throughout the Basin and Great Plains. This relationship is to be expected (Woollard, 1959), and the larger topographic features are associated with larger local positive free air anomalies.

The area of the Capulin free air gravity map (figure 8) is covered by volcanic rocks of three sequences - Raton, Clayton, and Capulin. The average free air anomaly over the map is about +55 mgal, with a large

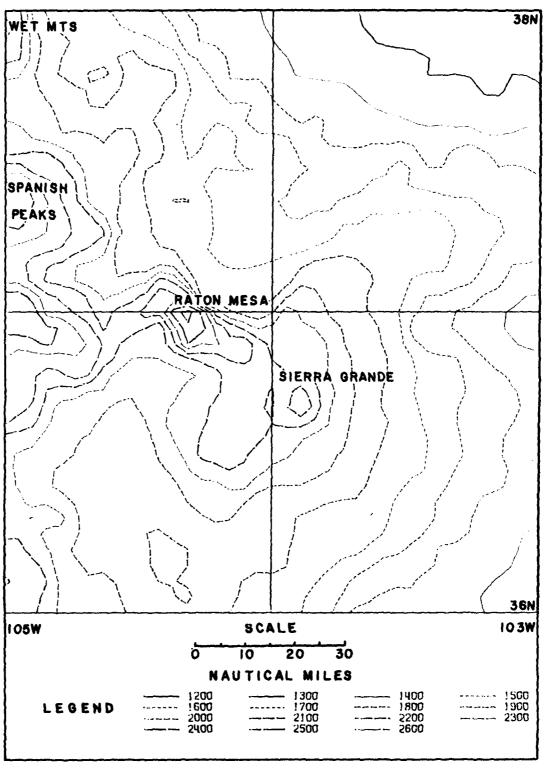


Figure 7. Regional topographic map of the study area. Contoured in metters, datum is sea level.

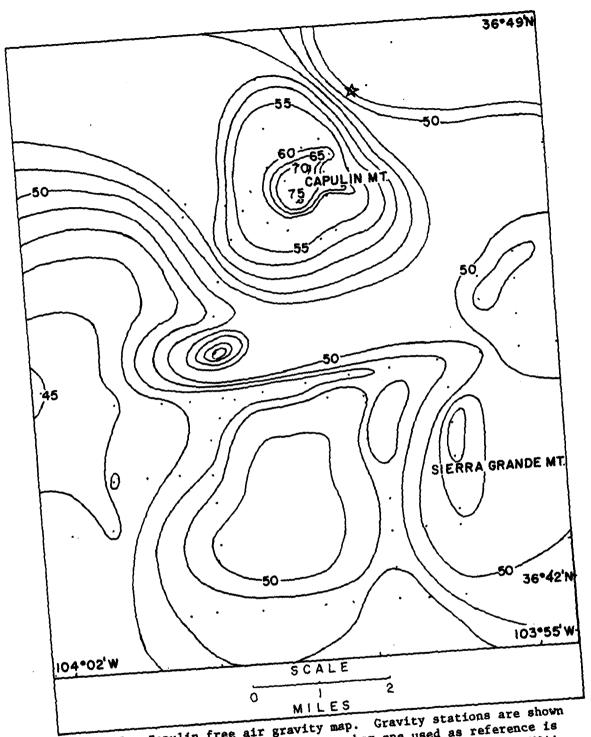
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Capulin free air gravity map. Gravity stations are shown by dots ('), and station number one used as reference is designated by a star. Station number one is also a NOAA Figure 8. gravity station, therefore, correlating the Capulin gravity data to the NOAA gravity data.

free air anomaly of +70 mgal associated with Capulin Mountain. This large free air anomaly at the mountain station (+70 mgal) is +15 mgal different than the regional free air value immediately around the mountain (+55 mgal). The 335 m cone comprising the mountain has approximately the same volume as a right cylinder 112 m high with the same radius of 500 m. The gravitational effect of such a cylinder, assuming a density of 2.67 g/cc, would be 12.4 mgal if the top of the cylinder is located at mountain peak elevation. This calculated value is not precisely equal to the 15 mgal anomaly observed associated with Capulin Mountain, but the method of calculation is only approximate. Agreement is close enough to allow the assumption that the gravitational effect of Capulin Mountain itself is 15 mgal.

Bouguer Gravity Data

The plot of average elevation against average Bouguer anomalies (from figure 6) indicates an inverse relationship between elevation and the Bouguer data (figure 9). The Bouguer anomaly is actually the negative gravity effect of the root increment below the compensation level. As the terrain elevation increases, the root increment below the compensation level also increases in thickness to maintain isostatic equilibrium. An increase in the thickness of the root below compensation level results in a more negative Bouguer anomaly. The Bouguer gravity maps (in hand drawn form - figure 10, and computer plotted form - figure 11) indicate this relationship between rising topography (figure 7) and a decreasing Bouguer equipotential surface. The slab assumed in making the Bouguer correction results in the computation of a Bouguer anomaly that is excessively negative in areas which are distinguished

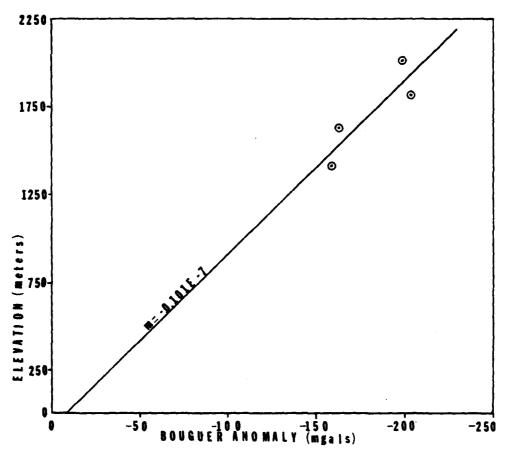


Figure 9. Plot of the averaged NOAA elevation data vs, the averaged NOAA Bouguer gravity data. The slope of the line, m, shows the inverse relationship of elevation and Bouguer gravity values.

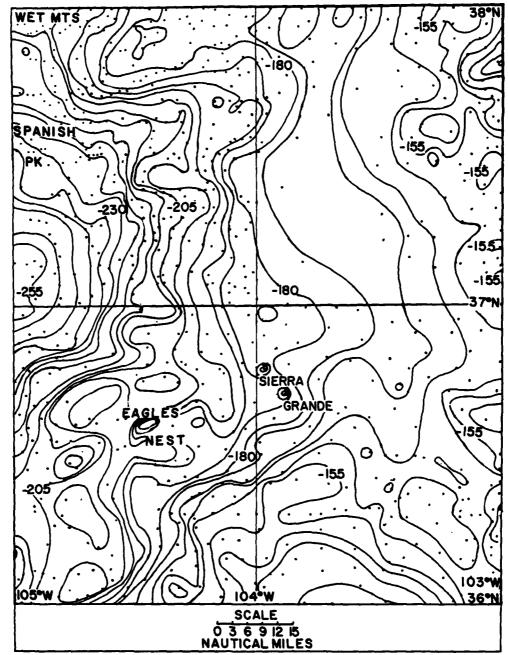


Figure 10. Regional Bouguer gravity map using NOAA gravity data. The dots represent the gravity stations. Datum is sea level, and a density of 2.67 g/cc was used in calculating the Bouguer correction.

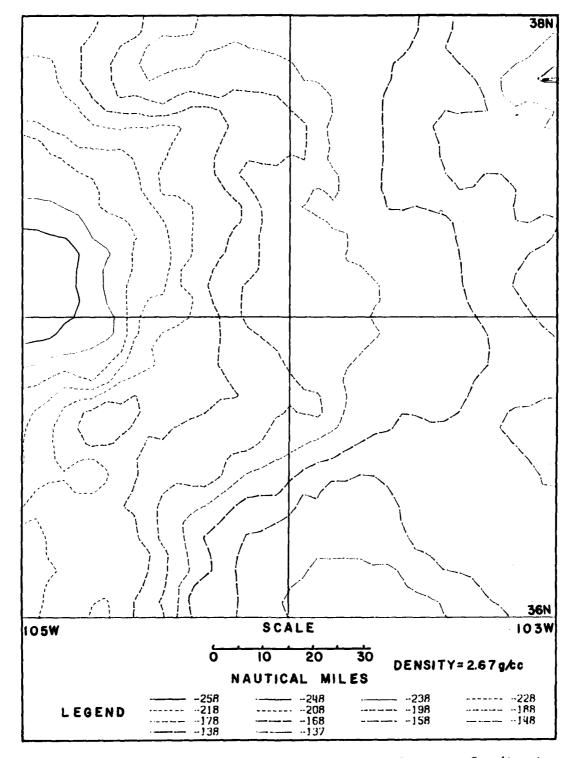


Figure 11. Regional Bouguer anomaly map of the study area. Gravity stations are the same as in figure 10. Contour interval is 10 mgal, datum is sea level.

by topographically isolated terrain, such as mountains. Sierra Grande Mountain is an example of this effect; however, Spanish Peaks and the Wet Mountains regions have sufficient areal extent to allow the slab assumption in the Bouguer calculation.

Eagles Nest Mountain has associated with it both a positive free air anomaly and positive relief with respect to the surrounding Bouguer surface on the Bouguer anomaly map. The positive free air anomaly is expected due to the near surface mass effect of the mountain. The presence of positive relief on the Bouguer equipotential gravity surface may indicate a near surface high density body, which would cause the Bouguer correction to be too small, or it may indicate the extension below datum of the igneous material forming the mountain. In either case, it appears that the density of the rocks forming Eagles Nest Mountain is greater than 2.67 g/cc.

The free air anomalies corrected for the Faye effect in the study area have already been shown to indicate isostatic adjustment, and the computation of isostatic anomalies directly from Bouguer anomalies can be used to add validity to this conclusion. The computation of the isostatic anomalies done here is based upon the Airy isostatic model with the density contrasts and formulas suggested by Woollard (1969). The slab thickness of the mass below the compensation level is determined by assuming that part of the negative Bouguer anomaly is caused by the presence of a slab of crustal material displacing heavier mantle material. Woollard (1969) determined that a crustal density of 2.92 g/cc is representative of the crustal rock at the compensation level (M-discontinuity)

and a density of 2.67 g/cc is representative of the upper crustal rock. The gravitational effect of the elevated surface is equal to the simple Bouguer correction computed using a density of 2.67. However, the effect of the surface slab upon a measured value on the surface is exaggerated as compared to the effect of the compensating root slab upon the same measurement as previously discussed. Therefore, if the region is in isostatic balance, the observed simple Bouguer anomaly is equal to the Bouguer gravity effect computed for the surface slab with average elevation in the area, minus the exaggeration effect of slab position. Previously, it was shown that this exaggeration effect is approximately equal to the observed free air anomaly. The isostatic anomalies for each one degree block in the study area have been calculated in this manner, and the values are shown in table 1.

These isostatic anomalies show the study area to be in near isostatic equilibrium, as was suggested in the free air anomaly discussion.

Qureshy (1962) studied the Bouguer data for all of Colorado and concluded that the eastern portion of the state, including the northern part of this report's study area, is in isostatic equilibrium. The slight positive bias in the results may be explained by either crustal thickening or increased crustal density (Woollard, 1969).

Table 1 includes calculations of crustal thickness using Shurbet's (1966) density model and formulas. The crust thins from west to east, and continues to thin east of the study area. Stewart and Pakiser (1962) calculated a seismic refraction profile using the Gnome explosion.

Their computed thickness is very close to the crustal thicknesses

Theoretical and actual gravity anomalies in the study area. All values averaged over one degree block. Table 1.

Area 10x10 N-W	Elevation h (meters)	Theoretical gravity due to the posi-	0bs. BA p=2.67 g/cc	Obs. FAA mgal	Isostatic anomaly FA-BA-gp	Compensation "root" H km	Crustal thickness T km
: :	. 6	tive mass g_p	mgal (1)	(1)		(3)	(4)
36-103	1636	182.8	-163.4	+19.9	+0.5	9.07	43.71
36-104	2010	224.6	-198.5	+26.2	+0.1	11.02	46.03
37-103	1412	157.8	-158.8	6.0-	+0.1	8.82	43.23
37-104	1818	203.1	-204.7	-1.2	+0.4	11.37	46.19
			-				

(1) from figure 6

(2) g_p = 2 $\pi\gamma\rho h$, where ρ = 2.67g/cc, and h = average terrain elevation

 $H = BA/2\pi\gamma\rho = BA/18.01$, where H is the thickness of the compensation root; p = 2.67 g/cc (3) from Shurbet (1966)

 $\tau = 33 + \mathrm{H} + \mathrm{h}$, where T is crustal thickness, H is root thickness, and h is terrain elevation (4) from Shurbet (1966)

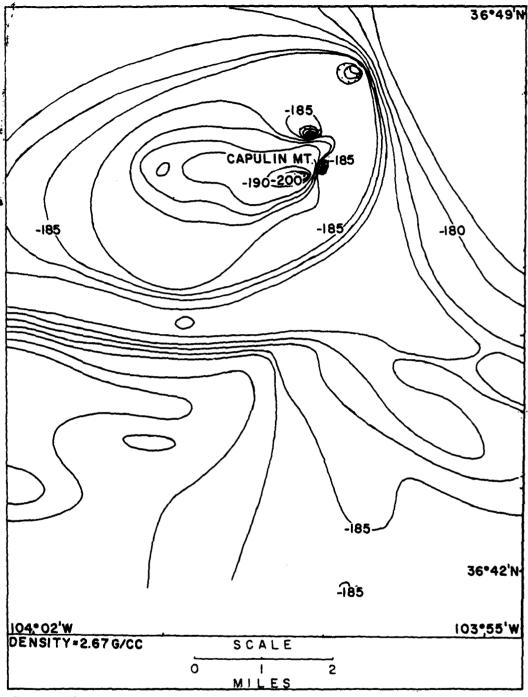
calculated using gravity data in this report.

In addition to the NOAA Bouguer gravity data, a separate Capulin Bouguer gravity map (figure 12) was prepared from data collected by the author. The Bouguer anomaly associated with Capulin Mountain is about -200 mgal, which is approximately 20 mgal more negative than the adjacent area. Since the mountain is a feature with little area, the slab assumed in the Bouguer correction has exaggerated the Bouguer anomaly. Because the mountain itself is 335 m high, a total of 37.4 mgal was subtracted from the measured value as a part of the total Bouguer correction for the height of the mountain. The actual gravitational effect of the mountain has previously been shown to be about 15 mgal. Therefore, the Bouguer correction applied to the mountain station was exaggerated by 22.4 mgal. Correction for this exaggeration shows the area beneath the mountain to be only slightly more negative than the surrounding area, and no additional mass deficiency exists beneath the mountain as compared to the surrounding area.

Since the study area, including the Capulin area, has been shown to be in isostatic balance, the Bouguer equipotential surface is expected to reflect variations in crustal thickness as well as local changes in crustal density. Therefore, the Bouguer anomalies can be used to calculate sedimentary thickness in the Basin. Calculations of this kind were carried out to produce a basement relief or structure map.

Raton Basin Basement Structure Map

In previous studies (Griggs, 1948; Baltz, 1965), subsurface structure in the Basin has been deduced from surface rocks and well log



Capulin Bouguer gravity map. Data stations are the same as shown in Figure 8. Density used for the Bouguer cor-Figure 12. rections is 2.67 g/cc.

data. A basement relief map based on structural data cannot be drawn because geologic mapping in the Basin is incomplete. However, Bouguer gravity data were used in this study to deduce crustal thicknesses and to determine basement structure (or relief). The basement surface computed is defined as the boundary between the less dense sedimentary rocks above and denser crystalline rocks below. Intrusive rocks that penetrate the strata are also classified as basement rocks because their presence causes effective thinning of the sedimentary column in the computation.

The best method for determining basement relief is based upon the realization that the Bouguer anomaly is actually the negative gravity effect of the slab or root which has displaced the heavier material of the sub-compensation level. In this study, it is assumed that the negative effect is locally affected by less dense sedimentary rocks displacing heavier basement rocks; therefore, the variation in the Bouguer gravity values can be interpreted in terms of variation in the thickness of the sedimentary rocks.

As a first step in constructing a basement relief map for the Basin it was necessary to ascertain the regional trend of the Bouguer gravity map. A second order polynomial surface was empirically determined to best represent the Bouguer gravity surface as discussed in Chapter II. Variation between this regional trend and the observed local Bouguer gravity values is assumed to be caused by differences in thickness and density of the sedimentary rocks.

The areal extent of the various lithologic units is shown in plate

1; Clark (1966) gives representative densities for the rock types present. Representative densities for the columnar sections (figure 3) for the Las Vegas and Raton Basins, respectively, were found to be 2.27 g/cc based on the densities reported by Clark (1966). The density contrast between the representative density for the Basin and the density used in the Bouguer correction (2.67 g/cc) is -0.4 g/cc. Estimations of lateral variations in the density contrast throughout the Basin were based on the following criteria: 1) sedimentary rock thinning near the Apishipa, Cimarron, and Sierra Grande Arches; 2) presence of volcanic rock in the southern portion of the Basin; 3) higher density intrusive rocks near Spanish Peaks and Wet Mountains regions; and 4) an average sedimentary rock density of 1.8 g/cc (producing a -0.87 g/cc density contrast) for the Great Plains (east of the Apishipa Arch - Sierra Grande Arch axes) as determined by Shurbet (1966). A representative calculation from the Bouguer gravity map (figure 10) along 37°N, 104°-105°W is included (figure 13).

At $37^{\circ}N$, $104.5^{\circ}W$ (figure 13), the Bouguer anomaly value, determined from the Bouguer gravity map, is -216 mgal. The regional Bouguer trend value (-210 mgal) was computed from a second order polynomial surface described in Chapter II. The residual gravity value (Δg) at $37^{\circ}N$, $104.5^{\circ}W$ is the difference between the Bouguer gravity value and regional trend value at that point (-210 + 216 mgal = +6 mgal). The density contrast value ($\Delta \rho$) at $37^{\circ}N$, $104.5^{\circ}W$, determined from the density contrast model, is -0.58 g/cc. The depth to the basement relief surface is found

Figure 13. Basement structure profile derived from NOAA Bouguer gravity data along 37°N, 104°-105°W.

> A. The regional Bouguer dip is determined by a smooth curve through the Bouguer anomaly equipotential surface.

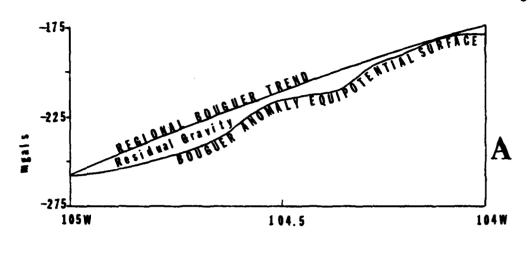
B. The difference of the regional Bouguer dip from the Bouguer anomaly equipotential surface in A is known as the

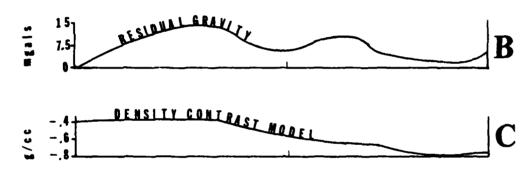
residual gravity (Δg).

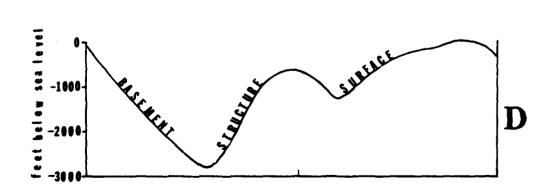
C. The density contrast model ($\Delta \rho$) is constructed from the known densities of the stratigraphic units in the Basin and subtracting those densities from the 2.67 g/cc density used in the Bouguer correction. Since the sediment density is 1.8 g/cc and it is assumed that the sediments are displaced crustal rock ($\rho = 2.67 \text{ g/cc}$), the gravity effect of this displacement is negative.

D. The depth to the basement structure surface (h) is calculated by using Δg from B and $\Delta \rho$ from C above in the

gravity mass effect formula: $h = \Delta g/2\pi\Delta\rho\gamma$.







by solving the gravitational effect formula for depth (h), using the Δg and $\Delta \rho$ values above: $h = \Delta g/2\pi\gamma\Delta\rho = -810$ feet (below sea level). This type of calculation was repeated for each of the 2,400 density contrasts in the Basin, and the resulting depths to basement structure were then contoured by the computer techniques described in Chapter II. The accuracy of the depths to the basement structure surface was independently verified using well log data (Baltz, 1965) and a measured section from the Raton Basin (McGehee, 1955).

In order to describe the correlation between the geology on plate 1 and the basement structure on plate 2, the same topographic names are used to refer to features on both maps. This practice is justified because the features named on the basement structure map all have assocated topographic features. In the area of the Wet Mountains and Spanish Peaks, the basement structure map shows that basement rock is exposed; that is, there are no sedimentary rocks present. The structurally high Apishipa Arch area on the basement relief map extends southeast of the Wet Mountains. The basement rises at about 1.5° to approximately 1,500 feet of vertical relief over a ten mile distance. The Apishipa Arch area is outlined by positive relief for approximately 50 miles on the basement structure map. The basement structure map shows the Apishipa Arch as the eastern boundary to the Basin.

The Delcarbon and La Veta syncline areas are shown on the basement structure map as depressions. The Delcarbon depression is above sea level and the La Veta depression is below sea level; both depressions

deepen to the south. The axis of the La Veta depression is the deepest depression on the basement structure map and defines the Raton Basin axis.

A positive relief feature on the basement relief map between the La Veta and Delcarbon depressions is also associated with intrusive rocks on the geologic structure map (plate 1). The shallow depth to known oil bearing formations of Cretaceous age (Baltz, 1965) make this anticline a favorable site for petroleum exploration. However, the heat from the intrusive activity may have destroyed any oil present, in a manner similar to the destruction of coal beds found elsewhere in the Basin and described in Chapter I.

The Cimarron Arch is an area of positive basement relief and clearly separates the Raton and Las Vegas Basins. It is asymmetric in shape, steeper on the north slope (4.5°) than the south slope (2.9°), and has approximately 3,500 feet of vertical relief. The Cimarron Arch area extends about fifty miles, from the west edge of the basement structure map to the Sierra Grande Arch area. The marked relief of this basement feature suggests the Arch has effectively prevented intertonguing of Las Vegas and Raton Basin's sedimentary rocks. As the Basins were downwarped and isostatic adjustment occurred to compensate for the accumulating sediments, the Cimarron Arch remained a positive structural feature. Although not a significant topographic or gravity feature today, the Arch was important in the development of the Basin.

The Sierra Grande Arch area is roughly defined by moderate relief on the basement structure map for approximately twenty miles east of

the Las Vegas Basin. Total positive relief increases about 1,000 feet over a twelve mile distance, producing a gentle 0.80 slope. The whole eastern sector of the basement structure map is a low relief positive area, rising slowly to the Apishipa and Sierra Grande Arch areas.

A previously unmapped and unnamed basement trough, which is in near alignment with a left lateral fault south of Spanish Peaks (plate 1), trends northeast from Raton Basin toward Apishipa Arch. The trough deepens southeast, towards the Raton Basin, and may be fault controlled. As a structurally low feature, the unnamed trough may contain significantly more sediments than areas surrounding it. The margins of the unnamed trough may prove favorable for future petroleum exploration. Presence of this trough is confirmed on the Raton Basin magnetic map.

Raton Basin Magnetic Map

The Raton Basin vertical magnetic data is tabulated and included in Appendix C. The magnetic data is shown in both hand drawn (figure 14) and computer plotted formats (plate 3). The computer plotted magnetic map is drawn at the same scale as the basement structure map for comparison. Zeitz et al., (1969) showed that the Colorado sedimentary rocks in the Basin are essentially nonmagnetic, the metamorphic and volcanic rocks are weakly to moderately magnetic, and the Precambrian quartz monzonite rocks are moderately to strongly magnetic. The Raton Basin magnetic maps actually represent the surface of the basement rocks in the Basin. Magnetic anomalies may also be associated with lateral changes in rock types, but there is no geologic evidence for basement

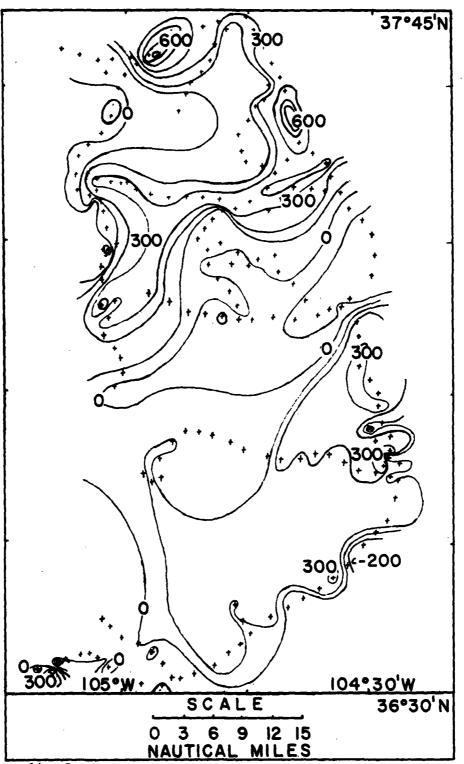


Figure 14. Raton Basin magnetic map. The crosses represent magnetic stations. Station number one is used as reference and is shown circled in the lower left corner of the map.

rock type changes in the Basin with the exception of those areas of known intrusive rocks.

Metasedimentary rocks are exposed near the Cimarron Arch and these rocks are near station number one, which was the magnetic survey datum. Magnetic measurements in the area of the metasedimentary rocks gave no indication that these rocks have higher magnetic susceptibilities than the sedimentary rocks in the Basin. Therefore, it is reasonable to assume that the rocks in New Mexico have the same magnetic susceptibilities that Zeitz et al., (1969) reported for Colorado rock types. Some lack of detail in the magnetic map in the southern portion of Raton Basin is the result of poor road access. All magnetic measurements were made along roads.

Structural depressions (plate 1) are associated with negative magnetic anomalies (plate 3) in the Basin. The La Veta and Delcarbon synclines both have distinct low magnetic anomalies associated with them; the unnamed trough detected on the basement structure map also has a distinct low magnetic anomaly associated with it. The magnetic anomaly associated with the unnamed trough makes it appear to deepen to the northeast, opposite the direction determined on the basement structure map. However, this apparent deepening probably reflects only the computer response to the lack of magnetic stations in the eastern part of the survey area.

Positive magnetic anomalies are associated with positive structural features. For example, the unnamed anticline between the La Veta and Delcarbon synclines has a significant positive magnetic anomaly

associated with it. The Cretaceous age sediments with known oil shows are only thinly buried near this anticline.

Capulin vertical magnetic data were also studied as an addition to the Raton Basin magnetic data. The Capulin data are tabulated and included in Appendix B with the Capulin gravity data. Unlike the sediment filled Raton Basin, the Capulin area is covered by volcanic rocks, which have significantly higher magnetic susceptibilities than the sedimentary rocks beneath. Some of the magnetic anomalies in the Cauplin area (figure 15) are associated with topographic features. For example, there is a magnetic anomaly associated with Capulin Mountain. However, a positive magnetic anomaly on the east slope of the mountain is not associated with a topographic feature. Free air gravity anomalies also mark this area, and this anomaly relationship may be an indication of lateral vents or dikes.

Variation in the magnetic properties of different volcanic rock types (Dobrin, 1976) would be helpful in identifying magnetic anomaly patterns in volcanic regions. Indeed, the largest positive magnetic anomaly is near Sierra Grande Mountain; smaller but equally distinctive positive magnetic anomalies are near Capulin Mountain. The unusually high magnetic susceptibility of the Sierra Grande sequence (indicated by its magnetic anomaly) may be diagnostic of these volcanic rocks and aid in determining their areal extent.

Capulin Local Seismic Study

In addition to gravity and magnetic measurements, an effort was made to determine natural seismicity in the Capulin area. It was

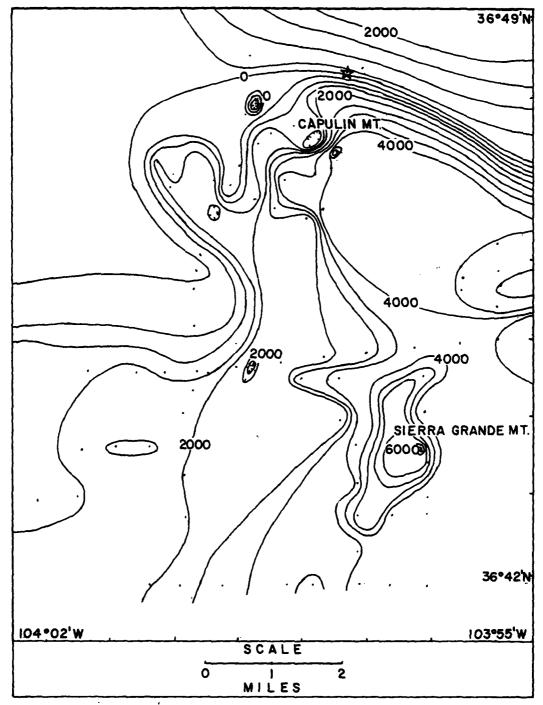


Figure 15. Capulin magnetic map. The dots represent magnetic stations. Station number one is used as reference and is represented by a star.

thought that some tectonic activity may be in progress. The results of the seismic survey in the Capulin area and all known seismic activity during the Twentieth Century are included in table 2. The epicenters of the earthquakes in table 2 are also shown on a map of northeastern New Mexico (figure 16).

A bar graph (figure 17) indicates the distribution of seismic activity in northeast New Mexico from 1900 to 1980. The lack of any significant recorded seismic activity during the first half of this century is due largely to the lack of recording equipment or populace to observe any shocks. Since 1962 the number of recorded events have increased due to the installation of permanent seismograph networks.

In the last five years seismic activity in New Mexico has declined (Sanford, 1981), yet this study suggests that the Capulin area has rather high seismic activity. The events recorded during this study were local in origin and small in magnitude, and were not recorded elsewhere

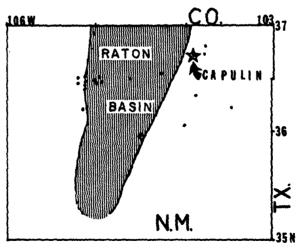


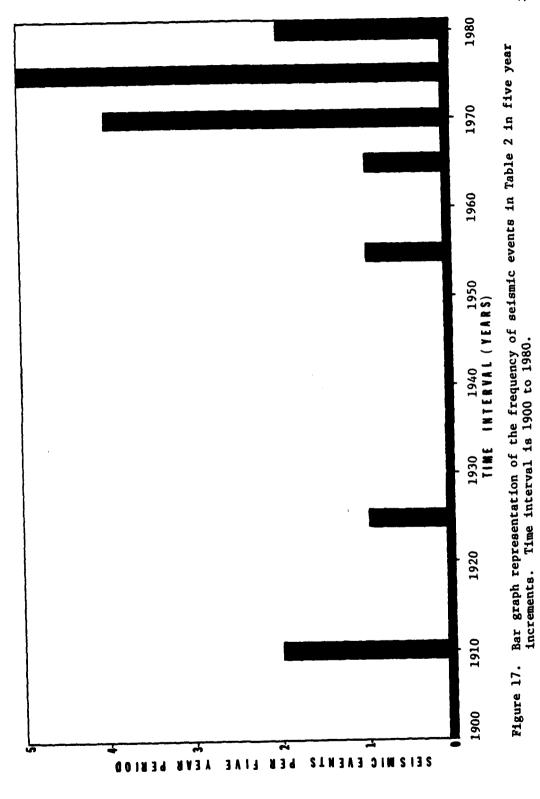
Figure 16. Geographic location of seismic events strong enough to be felt by local population in northeast New Mexico from 1900 to 1980.

Table 2. Recorded Earthquake Activity in the Study area of New Mexico 1900-1981.

SOURCE	DATE YR-MO-DAY	TIME GMT	LAT o _N	LONG OW	INTENSITY modified Mercali or local magnitude
1	06-04-19	2100	_	_	v
1	06-04-20	1300	-	-	VI
2	24-08-13	0423	36.0	104.5	V
2	52-08-13	2042	36.5	105.0	Ÿ
2 2	63-06-06	0806	36.6	104.4	2.7
2 2	66-09-24	0734	36,4	105.1	2.7
2	66-09-24	0827	36.4	105.1	2.4
2 2	66-09-25	1011	36.3	105.1	2.8
2	66-09-25	1223	36,5	105.1	2,8
2 2	72-02-20	2310	36.4	104.9	1.5
2	72-02-20	2323	36.4	104.9	2.2
	75-05-16	0138	36.9	105.0	1.5
2 2	75-05-16	0726	36.5	104.7	1.9
2	75-06-21	0542	36.1	104.0	2.0
3	79-09-19	0539	36.4	103.7	1.8
3	80-09-12	2138	36.5	105.1	2,3
4	81-07-24	1613		_	
4	81-07-24	1815			
4	81-08-27	0220			
4	81-08-27	0238			
4	81-07-27	1554			
4	81-07-27	1738			
4	81-07-30	1824			
4	81-08-25	2034			
4	81-08-27	2041			
4	81-09-01	2124			
4	81-09-01	2254			
4	81-09-09	1724			
4	81-09-16	0336			

Source Data:

- 1. Anonymous, The Raton Range (April 21, 1906) shock felt at Folsom, NM
- 2. Sanford et al., (1981)
- 3. Sanford (1981)
- 4. Capulin data (this report) epicenters assumed at or near Capulin, NM. No intensities were calculated, but the sensitivity of the equipment would indicate the relative intensities to be much smaller than previously recorded in the area.



in the state (J.J. Wolff, personal communication, 1981). Some of the events recorded may have been local mining blasts or sonic booms, although sonic booms are usually easily distinguished (figure 18).

The presence of seismicity near Capulin suggests this area to be either affected by dormant vulcanism, or by tectonic adjustment along faults in the area. Sanford et al., (1981) suggest that seismic activity in northeast New Mexico may be due to an eastern extension of the Jemez Lineament in western New Mexico. Continued monitoring of the Capulin seismicity would be necessary to determine the source and extent of the local seismic activity.

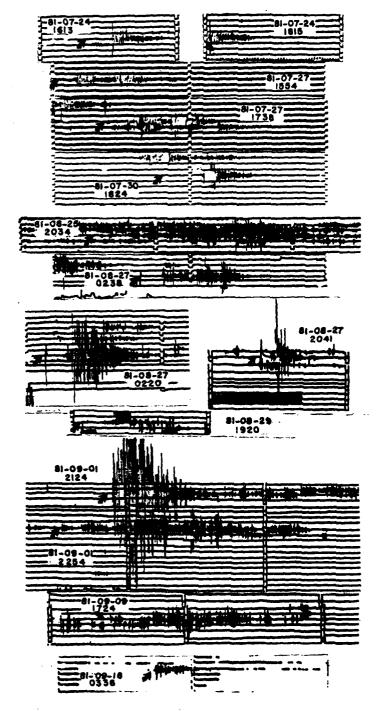


Figure 18. Seismograms from the Capulin seismic survey. The day-time groups correspond to Table 2.

CHAPTER IV

CONCLUSIONS

The geophysical measurements and the interpretation reported in this thesis meet the objectives set out in Chapter I. Furthermore, the interpretation of the data of these geophysical surveys lead to the following results:

- 1. The free air and Bouguer gravity data both indicate the study area is in near perfect isostatic adjustment. Correlation between this report and studies in adjacent areas indicate the Colorado northeast New Mexico west Texas region is in isostatic balance.
- 2. A large free air gravity anomaly in the southern portion of the study area is the result of local crustal loading by Quaternary volcanic activity. The thickness of the volcanic rock is estimated to 180 m based on the gravitational effect represented in the free air anomaly.
- 3. The basement relief map constructed from Bouguer gravity data has a high correlation with known geologic structures.
- 4. The Raton magnetic map generally represents the same basement surface as that constructed from Bouguer gravity data.
- 5. The Capulin area and adjacent High Plains are seismically active. Further investigation of the local seismicity is important.
- 6. The basement structure map and Raton magnetic map indicate areas of potential benefit in future petroleum exploration. The high heat flow in the Basin may have economic importance. If the heat flow is found to be a result of dormant vulcanism, seismic studies, as in the Capulin area, may be helpful in determining the extent of the thermal field.

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APPENDIX A

NOAA GRAVITY AND ELEVATION DATA

Explanation of Table Headings

Lat Northern latitude of station in degrees.

Long Western longitude of station in degrees.

Elev Station elevation in meters above sea level.

FAA Free air anomaly, calculated from a datum of sea level. The free air correction (FAC) is 0.3085 mgal/m and is applied to the observed gravity reading (OGR): FAA = OGR + FAC.

BA Bouguer anomaly, calculated using a datum of sea level, station elevation, and a density of 2.67 g/cc. The Bouguer correction (BC) is $2\pi\gamma\rho h$, or 0.0837 mgal/m, and is applied to the FAA: BA = FAA - BC.

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36.844 103.327 36.91? 103.227	1632.300 1457.900	10.600 -0.300	-171.500 -167.400
30.413 103.365	1488.200 1352.500	-5.30C	-167.300 -171.800
35.712 103.234 35.714 103.034	1352.500	14.300	-159.200 -153.300
35.714 103.034 35.735 123.414	1474.400	3.500 2.400	-175.100
30.533 103.360	1053.600	19.1čč	-165.7CC
30.007 103.444	1723.300	25.305	-166.400
36.757 103.307 20.735 103.395 35.833 103.017	1528.500	10.400 15.600	-155.000 -17:500
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30.948 103.408 30.917 103.090	1537.000	-2.100	-174.166 -160.700
36.969 103.195	1364.200 1404.500 1481.100	-8.100 -7.700	-164.500
36.772 103.345	1431.100	~3.00C	-108.706
36.750 103.138 36.502 103.279	1325.600	7.900	-162.700
25.534 103.121	1477-600	1 J. 600 6. 500	-162.200 -158.346
30.000 103.324	1925.700	.5.500	-163.200 -167.700
30.650 103.469 30.050 103.181	1741.100	25.900	-167.700
35.612 153.072	1489.000	18.500	-121.700 -147.900
35.517 103.072 35.517 103.487	1741.900	25.900 20.000 18.500 19.900 19.900	-106.200
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37.777	1282.700	-y.400	-152.900
37.792 103.547 37.996 103.133	1314.000 1217.400	-13.000 -23.000	~100. ≥€€
37.996 103.133 27.851 103.090	1217.400	-23.000	-156.200
37.775 103.100 37.775 103.100	1.65.300 1264.000	-8.000 -12.400	~150.200 ~153.900
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37.939 103.055 37.939 103.200	1230.500	-13.200 -13.500	~151.000 ~154.000
37.939 103.200 37.881 103.390	1292.403 1265.500 1264.300	−8.500	-150.100
37.881 103.390 37.830 103.200	1264.000	-12.300	-150.100 -153.700
	1280.500	5.6 2.700 -0.100	-143.100 -139.500
37.118 1.3.100 37.148 103.148	1275.000	~ō.166	-144.UCC
37.7+d 103.060	1330.700	3.200	-143.4CC
37.090 103.120 37.091 103.042	1313.400	-3.800 1.500	-150.800 -150.000
37.677 103.093	1325.400 1325.300	-2.400	-150.000
37.077 103.043 37.034 133.040	1325.300	0.760	-147.300
37.611 103.625 37.536 103.002	1+14+300	4.600	-140.000
37.529 103.065	1449.300 1485.900	9.000 9.800	-156.000 -157.300
37.253 103.067	1400-900	7.500	-155.700
37.782 103.137	1289.500	0.400	-137.700 -130.100
31.803 103.025 37.848 103.225	1290.300 1274.660	-10.600	-130.100 -153.100
37.003 lu3.367	1274.000	~!). 400	-143.000
3/-/38 133-26/	1314.900	~1 3.630	-157.8CC
37.703 103.009 37.788 103.398 37.995 103.233	1352.400 1258.800	-15.400	-149.000 -156.200
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APPENDIX B

CAPULIN GRAVITY AND MAGNETIC DATA

Explanation of Table Headings

Numerical listing of gravity and magnetic station iden-Station tification. Stations with apostrophe (') are addition-

al magnetic stations.

Northern latitude of station in degrees. Lat

Western longitude of station in degrees. Long

Free air anomaly (FAA) calculated from a datum of sea Free Air An. level. The free air correction (FAC) is 0.3085 mgal/m and is applied to the observed gravity reading (OGR):

FAA = OGR + FAC.

Bouguer anomaly calculated using a datum of sea level, Bouguer An. station elevation, and a density of 2.67 g/cc. The Bouguer correction (BC) is 2πγρh, or 0.0837 mgal/m

and is applied to the FAA: BA = FAA - BC.

Relative vertical magnetic reading in gammas. Datum is Rel. Mag.

station one.

Station	Lat	Long	Free Air An.	Bouguer An.	Rel. Mag.
1	36.805	103.956	49.0	-186,2	0
2	36.806	103.951	48.8	-183.1	695
2'	36.804	103,960			1710
3	36.800	103.965	57.89	-184.5	1395
3'	36.796	103.963	2		2070
4	36.792	103,963	64.0	-179.5	925
41	36.790	103.965			850
5	36.794	103.961	59.25	-185.8	1340
51	36.792	103.966			1790
6	36.790	103.978	60.0	-187.1	1950
7	36.796	103.982	58.5	-186.6	275
71	36.798	103.979	••••		-1200
8	36.809	103.980	53.0	-183.5	750
9	36.785	103.983	57.85	-189.3	75
10	36.782	103.969	57.06	-188.7	145
11	36.776	103.973	54.4	-187.9	600
12	36.767	103.969	50.6	-186.7	275
12'	36.755	103.978	•		-150
13	36.759	103.979	49.3	-186.6	200
14	36.752	103.978	54.8	-181.1	450
15	36.743	103.978	47.63	-186.3	1125
16	36.743	103.992	47.8	-186.2	1500
17	36.744	103.984	47.7	-186.2	2425
17'	36.735	103.979			1250
18	36.744	103.973	47.9	-186.1	2125
19	36.746	103.962	47.5	-184.5	2350
20	36.746	103.951	49.05	-182.6	2775
21	36.748	103.938	49.1	-181.6	3625
22	36.748	103.927	50.05	-181.2	3525
23	36.749	103.917	49.9	-181.3	3500
24	36.779	103.979	58.45	-189.4	1375
25	36.784	103.964	59.3	-204.2	3500
26	36.783	103.971	74.6	-195.0	3800
27	36.756	103.917	49.85	-180.2	4650
28	36.759	103.927	48.7	-180.6	4600
28'	36.762	103.926			4525
29	36.764	103.927	49.1	-180.5	4500
30	36.764	103.917	49.5	-179.7	4750
30'	36./61	103.917			5000
31	36.769	103.917	48.1	-179.3	4700
32	36.742	103.936	49.9	-184.2	4920
33	36.734	103.936	52.5	-182.5	3750
34	36.727	103.935	52.3	-183.2	6050
35	36.723	103.935	51.6	-184.5	5500
36	36.716	103.935	51.3	~185.5	3650
37	36.713	103.935	50.85	-185.7	3860
38	36.713	103.944	50.3	-184.7	3950
39	36.713	103.950	48.8	-185.0	46 xi

Station	Lat	Long	Free Air An.	Bouguer An.	Rel, Mag.
40	36.705	103.953	48.0	-185.4	3300
41	36.698	103.952	47,75	-185.0	3600
42	36.698	103,945	47.8	-185,4	4000
43	36.698	103.934	49.3	-185.1	3600
44	36.741	103.968	49.1	-184.5	3450
45	36.741	103.958	48.9	-183.3	3600
46	36.732	103.958	48.0	-184.5	2450
47	36.727	103.953	48.0	-184.4	3750
48	36.719	103.953	48.4	-184.7	3500
49	36.689	103.965	48.4	-185.3	2850
50	36.689	103.972	48.6	-185.6	3200
51	36.689	103.988	48.2	-186.4	2700
52	36.689	104.023	48.1	-187.7	2000
53	36.689	103.999	48.3	-186.5	2350
54	36.712	103.998	52.1	-186.3	1750
55	36.721	103.998	49.3	-187.7	2000
56	36.727	103.999	48.5	-186.9	1650
57	36.734	103.998	47.9	-188.6	1900
58	36.739	103.996	47.8	-187.3	1900
60	36.743	104.003	47.3	-186.1	1375
61	36.743	104.011	45.9	-186.9	1700
62	36.743	104.021	45.85	-187.7	1275
63	36.746	104.042	44.85	-188.3	1375
64	36.727	104.007	47.7	-186.7	1225
65	36.726	104.016	46.3	~187.5	1400
66	36.726	104.025	47.7	-187.5	1725
67	36.708	104.025	45.7	-188.9	1425
68	36.708	104.036	46.8	-189.8	1325
- 69	36.787	104.005	52.2	-188.2	1775
70	36.784	104.000	52.55	-187.6	1375
71	36.781	104.001	51.4	-188.1	1500
72	36.828	103.990	53.9	-189.8	1725
73	36.789	103.958	54.4	-187.2	5075
74	36.777	103.962	54.4	-186.0	4000
75	36.777	103.974	57.9	-188.3	2675
76	36.792	103.966	54.4	-185.7	2125
77	36.786	103.960	68.1	-180.6	3900
78	36.789	103.964	58.1	-188.8	2975
79	36.815	103.945	49.5	-178.4	2250

APPENDIX C

RATON BASIN MAGNETIC DATA

Explanation of Table Headings

Lat Northern latitude of station in degrees.

Long Western longitude of station in degrees.

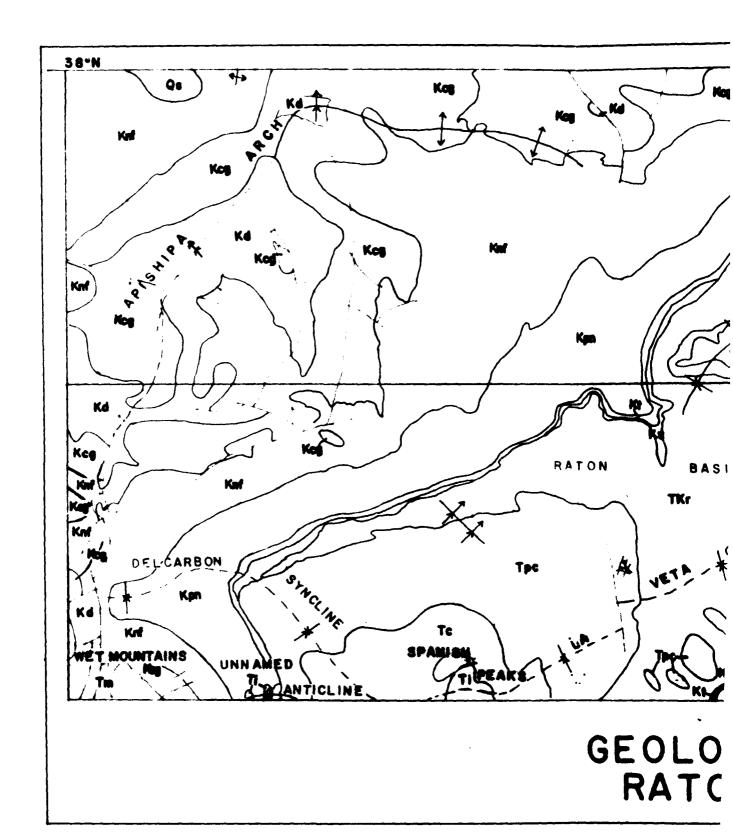
MAG Relative vertical magnetic reading in gammas. Station one is

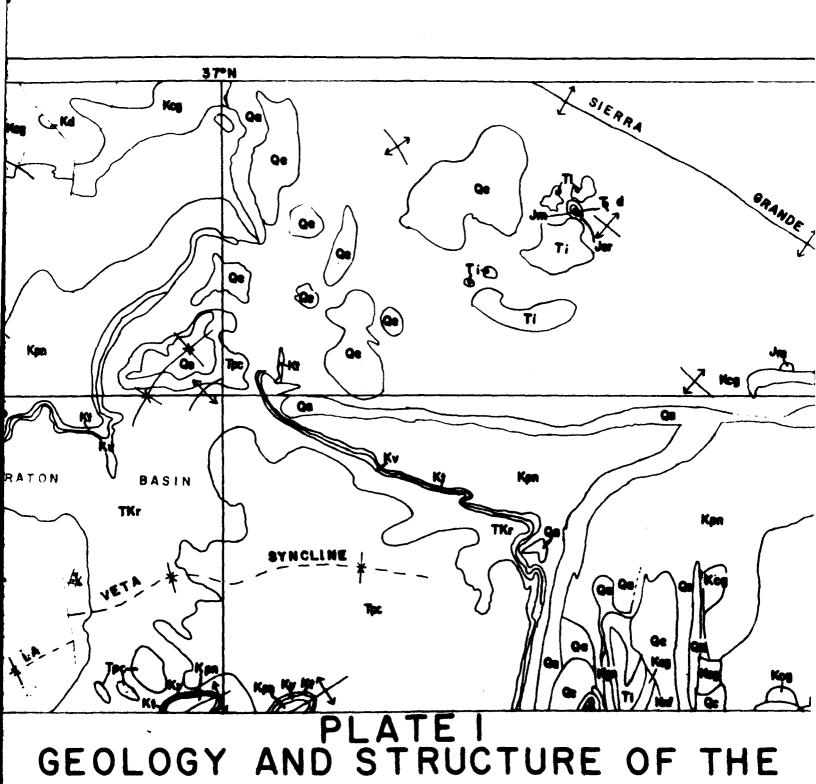
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	105.17	5:.7_0	35.530	135.110	-111040
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යට වේ 🕼	105.030	<u>, −5 • 2 2(/</u>	20.576	17.676	
30.235 20.249	100.000	-114.400	ઉ ⊍• ા+િ	主 リン・レゴレ	24.400
2 3 3 5 6	lub.Clu	-7.520	36.3 <u>.</u> 0	エンチャンプロ	-los.220
- CO •		-24. <u>25</u> 3	30.32.	104.464	-114.000
ခုံပါမှာဖိပြ	1040700				142.010
3 ⊙ •♂3∨	144.244	1∠5.→ÇÇ	30.32F	1-4-246	144.010
ك 3 و ب	104.450	10%50	35.700	· J+ · 4/4	13.00C
33.100	104.490	-1.510	30.750	134.510	-1740026
Jo .725	و د د ۱۰۵		35.11.	シリー・シャレ	1.01.
		-201.040		194.066	1,7.000
30.096	iv4.570		50.070	754.000	-7.000
<u> </u>	104.030	-33.02 <u>0</u>	20.040		
35.030	164.73	105.580	35.00€	104.740	57.540
الأسان والأسان	.04.750	11.5.0	30.040	104.176	المان واشتاذ الأ
30.023	134.735	76.780	30.360	114.15C	llwarit
30.4334	104.130	19.280 120.230		l. + . clu	133.3
20 • ⊃0 i	104.75	* 50 • 5 3 V	200340	107.67.6	
يا د ز. ن ت	104.850	-29.150 -54.740	30.530	194.850	-54.150
30.526	104.87.	ータしょ7しご	さいる シンコ		-20.44.0
30.07.	104.940	-126.426	jn.55.	134.710	しゅう きゅ
			Ĵα.∪Ι∪		ت به م ° ر ⊷
ئۈر• باڭ	104.970	~დე.აგ გ ნ		10000	12.240
30 .5 By	105.040	-72.480	30.30U	134.400	1-0-70
ဖြင့် စော်ခဲ့	134.497	2 9 7.200	30.J	1 4	
23.035	104.050	1/5.190	35-775	154.000	132.1-0
33.092	134.013	210.170	30.000	134.640	150.314
		11.4		134.710	45.000
يان وي دو	194.732	212.350	30.340		
30.3333	134.750	71.470	36.910	104.700	56.720
33.736	164.322	75.130	うい・ソンじ	エリサ・ジャレ	29.324
رَوْ بِالْحِيْرِ عُوْدُ	104.550	55.210	36. 7.45	1 34.360	lioeció
		33.41	36.800	1,4.920	91.1.0
20.007	164.930	33.410		174.900	13.040
30.792	194.340	112.510	300 200		
30.233	1040/	331.756	46.09	194.40%	19.900
ياران و ي	10+.400	5301070	20.940	134.776	259.136
	104.530	-80.130	30.455	134.400	233.000
50.424			30.700	134.470	312.110
35.755	134 . +]	39.100	20 • 7 4 0		
33.79.7	104.493	224.100	37.030	14.206	4340676
.7	104.57.	130 . O C C	37.570	よりみょうこじ	346.44
31.112	134.530	207.410	37.130	134.530	2140666
7.4		-171.090	37.150	. 34	ーチン・フェレ
37.103	الكوائر والمراجع المراجع				
51.L+U	104.290	-125.780	21.12U	1,14.056	-132.310
51.137	104.770	-13.750	37.13.	1140146	-09.454
37.120	104.770	-62.34C	31.136	104.506	-111.500
3 4 4 7	104.030	_13145	37.130	134.670	81.000
37.133	104.03	-19.330 151.910	5/.1/	134.456	353.476
27.150	104.900	737.27	34.1		
27.152	105.0+0	0 ن 7 - 74	37.13C	يونيا بالمائزان	424.0.0
31.170	105.050	135.770	37.210	104.070	94.004
31.190	105.050	13.220	31.22	104.7.0	40.260
31.21.	104.150	135./70 13.250 -22.540	37.240	134.710	-30 ·420
51.00		32.074	77.5	14.190	16.420
14.4234	I (+ •51 -	24 + 7 / %	37.230 37.230		
37.253 31.203 27.273	104.576	-120.460	37.55	174.2-6	-1+7-130
27.270	104.500	-170.000	37 90	104.5.0	-150.500
31.310	104.550	-170.030	31.330	104.576	14.300
7		197.120	37.30.	134.000	بال في و و و و
_{	104.29	224			457.750
يال4. إد	104.520	336.240	37.420	104.000	
47.500	164.375	321.326	27. 270	104.700	200.920
21.290	104.72	439.37.	37.320	134.750	301.120
57.510	- そろるこうほど	300 . L 30	57. 3.70	1040626	3320-40
	10 T # 1 2 4		37 . 57	134.576	303.326
31.320	104.340	303.290	37.52G		303.320
11.0	ショチャラニン	i kw.⊾ow	٧٤٥٠) د	エッテ・デンシ	-73.4-6
3/ 3/30	104.980	344.676	21.25C	165.666	337.576
55 55	1.5	272.470	37.400	135.000	302,150
47.7	105.010 105.000		27. 220	135.350	442.000
37.329	Tanadaa	၁၉ 7 . ၁၀၆			
د <i>ک</i> فی ۱ د	1000g/c	172.280	3 1 . 3 4 9	135.039	11.046
31.311	104.710	292 . 046	27.320	104.090	اداء∠پ3
31.325 31.345 31.345	164.67	11.169	ت€د • 7 ف	10+0530	∡ċ5•००७
31.333	104.000	235.640	37.440	Ĩu4.660	007. : JU
		237	27.470	114.750	494.4.0
↑₹•+7 <u>५</u>	104.07	027.			
21.02-0	1.4.7	يرود ، آور	37.34.	1,4.11	342.124
37.57	104.700	الأطلاء تارات	21.370	124.746	250.346
-					

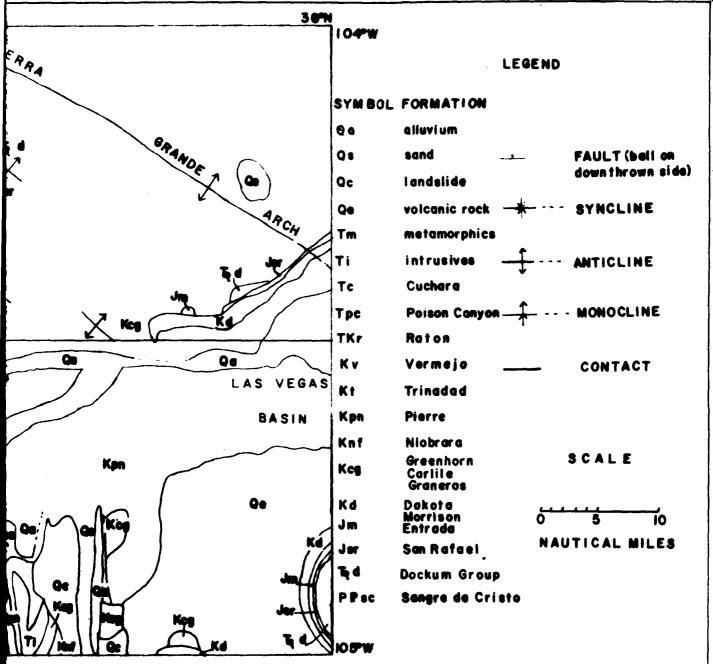
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		440	<u>_</u>		<u> </u>
37.520	1.4.779	247.110	37.01.	155.46	327.5.6
<i>□ [• > □ □</i>	ڭۇئى بەن ∡	306000	37.000	1 14 + 60 C	430.00.0
3/ 02/6	14.444	009.150	27.700	بالم لأ و بدر م	115.510
57.50	194.96	620•37¢	57.120	نَا يُرَبُّ مَا ﴿ فِي مِنْهِ	3,1,020
<i>⊃لای</i> اید	105.040	266.556	21.676	105.626	100.610
3 7 . 000	102.000	1+7 10	37.626	<u> </u>	46.400
37.020 37.2+0	بال، قبيا	-37.j	31.210 37.240	بادياء ذرا	~23.41 0
	102.0+0	-55.020	3/0240	1,700,00	13-126
37.270	105.040	42.700	37.290	135.050	بانائدهدي
ر2د.)د	102.07.	- 37 . / 10	21.200	المراجية ودا	14.466
<u>2,7</u> .20€	105.100	57.75Q	3/.410	102000	្តភ្ទឹកក្នុងC
37.420	105.023	=30 €320 =133 03	37.	402.650	-194.20g
3/.480	100.040	-133.91	3 7. 500	1-5-420	-23.000
21.543 47.548	103.010	40 42 FU	27. 240	And the Property of	اجلاءييت
37 • 340 27 • 340	104.990	134.500	21.540	105.000	22.370
27 Ex.3	105.130	22.100	37.520	1-2-100	30 + <u>2</u> + €
37.500 37.5yı	104.630	130.436	37.520 37.576	145.656	.60.776
21.243	1.4.830	222.930	17 5/0	1.4.000	299.600
57.518	104.370	177.350	37.020 31.490	1.4.690	259.810 20.400
3/ 460	167.356	64.590	37.510	104.710	103.020
37.+60 37.+60	104.72	J27.3+6	37.130	134.11.	2030020
21.110	164.790	ان و اد	31.190	lu+.uli	47.446
37.220	104.030	-31.150	37.240 37.270	434016	5.010
31.200	104.853	11.170	37.276	134.340	3.140
37 .I JÕ	104.330	73.250	51.20.	204.000	30.11
21.330	10+.730	144.416	21.300	1,4.760	104.726
57.59	104.775	Ījājātu		1 14.736	157.200
31.453	104.750	151.310	31.410	134.740	246.446
37.130	104.010	-162.796	27.10.	104.030	-174.550
37 - 19 -	104-000	-47 × 741.			





RATON BASIN OF N.M. AND COLO.



OF THE COLO.

38°N UNNAMED TROUGH 1500 -3000 DELCARBON SYNCLINE

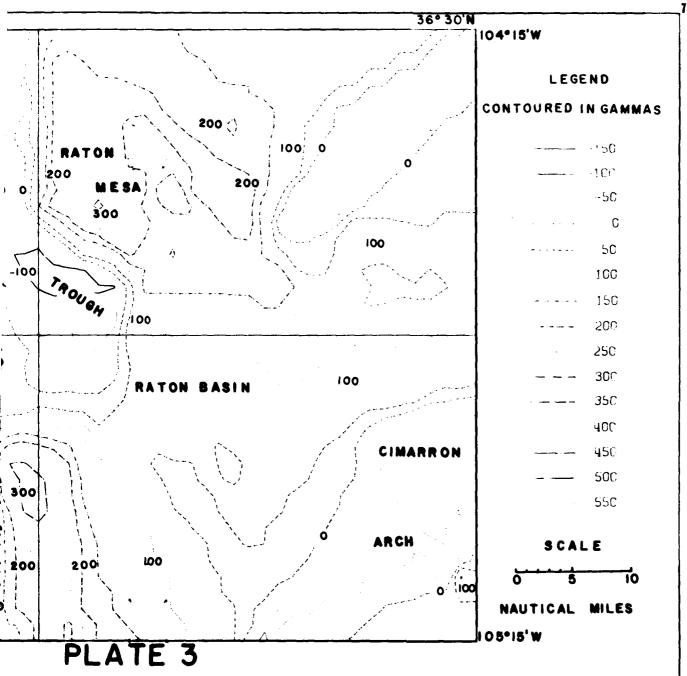
> BASE RA'

RATON BASIN OF N.M. AND COLO.

36°N enanor LEGEND DATUM IS SEALEVEL BASIN SCALE -1500 NAUTICAL MILES DR THE ('OLO.

GROUND MAGNETIC MAP RATON BASIN OF N.M. A

DATUM IS STATION ONE IN FIGURE



MAGNETIC MAP FOR THE SIN OF N.M. AND COLO.

TUM IS STATION ONE IN FIGURE 14

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